

Signatures of terminal alkene oxidation in airborne formaldehyde measurements during TexAQS 2000

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[1] Airborne formaldehyde (CH₂O) measurements were made by tunable diode laser absorption spectroscopy (TDLAS) at high time resolution (1 and 10 s) and precision (± 400 and ± 120 parts per trillion by volume (pptv) (2σ), respectively) during the Texas Air Quality Study (TexAQS) 2000. Measurement accuracy was corroborated by in-flight calibrations and zeros and by overflight comparison with a ground-based differential optical absorption spectroscopy (DOAS) system. Throughout the campaign, the highest levels of CH₂O precursors and volatile organic compound (VOC) reactivity were measured in petrochemical plumes. Correspondingly, CH₂O and ozone production was greatly enhanced in petrochemical plumes compared with plumes dominated by power plant and mobile source emissions. The photochemistry of several isolated petrochemical facility plumes was accurately modeled using three nonmethane hydrocarbons (NMHCs) (ethene (C₂H₄), propene (C₃H₆) (both anthropogenic), and isoprene (C₅H₈) (biogenic)) and was in accord with standard hydroxyl radical (OH)-initiated chemistry. Measurement-inferred facility emissions of ethene and propene were far larger than reported by inventories. Substantial direct CH₂O emissions were not detected from petrochemical facilities. The rapid production of CH₂O and ozone observed in a highly polluted plume (30+ parts per billion by volume (ppbv) CH₂O and 200+ ppbv ozone) originating over Houston was well replicated by a model employing only two NMHCs, ethene and propene. *INDEX TERMS:* 0345 Atmospheric Composition and Structure: Pollution—urban and regional (0305); 0365 Atmospheric Composition and Structure: Troposphere—composition and chemistry; *KEYWORDS:* formaldehyde, tropospheric oxidation, urban pollution, alkene oxidation, petrochemical facility emissions

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1. Introduction

[2] Gas-phase formaldehyde (CH₂O) is a ubiquitous component of the troposphere and is typically the most abundant carbonyl compound found in the boundary layer atmosphere. CH₂O is largely produced in situ, as an intermediate

in the oxidation of a large array of volatile organic compounds (VOCs). However, only a small number of VOCs are both abundant and efficient CH₂O producers. In the polluted boundary layer the terminal alkenes are generally the most important CH₂O precursors, particularly isoprene (C₅H₈), ethene (C₂H₄), and propene (C₃H₆) [Dodge, 1990; Lee et al., 1998b; Goldan et al., 2000]. In many cases, the important CH₂O precursors also contribute substantially to total hydrocarbon hydroxyl radical (OH) reactivity ($\Sigma[k_x\text{OH} * [\text{VOC}_x]]$), and thereby total VOC reactivity. Here k_x is the bimolecular rate constant for reaction of OH with each VOC. Due to the heterogeneous spatial distribution of CH₂O precursors, and the relatively short lifetime of CH₂O (2–4 hours in the sunlit troposphere), continental boundary layer CH₂O concentrations span nearly 2 orders of magnitude [Grosjean, 1991; Fried et al., 1997; Lee et al., 1998b], providing a sensitive indicator of recent VOC oxidation.

[3] CH₂O is also emitted directly into the atmosphere by incomplete fossil fuel combustion [Sigsby et al., 1987; Altshuler, 1993; Anderson et al., 1996], biomass burning

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[Lee *et al.*, 1998a; Yokelson *et al.*, 1999], industrial processes, and vegetative emissions [Carlier *et al.*, 1986]. However, primary CH₂O emissions are typically small relative to photochemical production. Using ambient measurements, Li *et al.* [1994] have estimated that the amount of CH₂O directly emitted from eastern North America is 16 times less than that produced by the oxidation of emitted anthropogenic and biogenic VOCs.

[4] Since VOC oxidation also drives tropospheric ozone production in the polluted boundary layer [Haagen-Smit, 1952; Chameides *et al.*, 1992; Sillman, 1999], CH₂O measurements are useful to deconvolving sources of ozone pollution, an issue central to urban and regional air quality. For instance, CH₂O measurements made during the 1995 Southern Oxidants Study (SOS 95), based in Nashville, TN, were used to corroborate the contribution of biogenic isoprene to regional ozone production [Lee *et al.*, 1998b]; and satellite measurements of CH₂O have been used to evaluate biogenic isoprene emissions over North America [Chance *et al.*, 2000; Palmer *et al.*, 2001]. To similar ends, the present study utilizes high time resolution airborne CH₂O measurements collected during the Texas Air Quality Study (TexAQS) 2000 to examine the role of various hydrocarbons in the photochemistry of the Houston, Texas area.

[5] During the 1990s, ozone levels in exceedance of National Ambient Air Quality Standards (NAAQS) were regularly reported for the greater Houston, TX metropolitan area. On average, only Los Angeles reported more intense episodes of ozone pollution during this period [Lin *et al.*, 2001]. Over 4 million people live in the greater Houston metropolitan area [U.S. Census Bureau, 2000], resulting in large mobile source emissions of VOCs and NO_x [TNRCC, 1996]. One of the largest electric utility power plants in the nation, the W.A. Parish facility, is located just outside of Houston (see Figure 1) (D. K. Nicks *et al.*, Fossil-fueled power plants as a source of atmospheric carbon monoxide, submitted to *Atmospheric Environment*, 2002). In addition, a large fraction of the U.S. olefin production capacity is found in the Houston–Galveston–Brazoria area, with an especially high density located along the Ship Channel inside Houston city limits (see Figure 1). Biogenic emissions are also a source of VOCs in and around Houston, although detailed vegetation surveys [Wiedinmyer *et al.*, 2001] indicate that these emissions should be substantially less than in areas such as Tennessee [Geron *et al.*, 1994; Goldan *et al.*, 2000].

[6] The TexAQS 2000 mission was conducted to elucidate the role of such emissions, along with meteorology, in Houston's intense photochemical pollution [Kleinman *et al.*, 2002]. The study was conducted from mid-August to mid-September 2000 and involved the combined efforts of the Texas Natural Resource Conservation Commission (TNRCC), the National Oceanic and Atmospheric Administration (NOAA), the National Center for Atmospheric Research (NCAR), Brookhaven National Laboratories (BNL), the University of Texas, Baylor University, and many other laboratories and universities. Multiple well-instrumented aircraft and ground sites were employed.

[7] The present study addresses four primary questions utilizing airborne CH₂O observations collected during TexAQS 2000: (1) What were the major sources of CH₂O (and

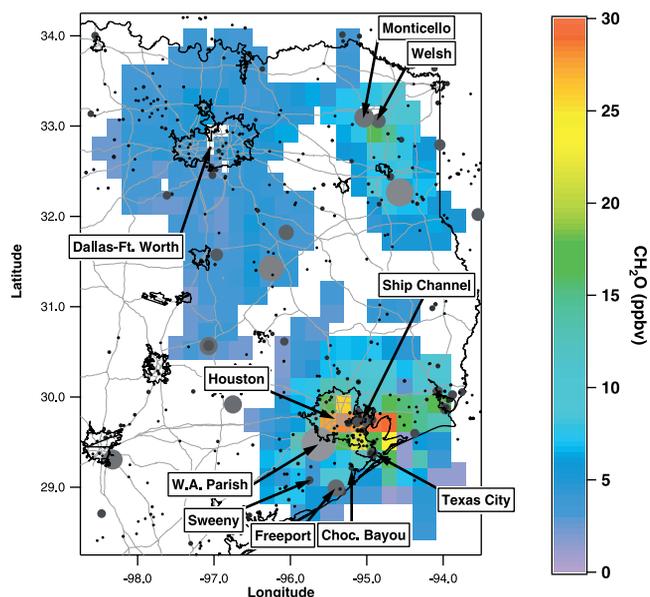


Figure 1. Regional distribution of CH₂O maximums measured below 1000 m flight altitude for 0.2° latitude by 0.2° longitude grid cells. The solid circles show the location of NO_x emitters and are sized according to reported NO_x emissions. Labels are provided for cities (Houston and Dallas–Fort Worth), utility power plants (W.A. Parish, Monticello, and Welsh), isolated petrochemical facilities (Sweeny, Freeport, and Chocolate Bayou), and areas of concentrated petrochemical production (Ship Channel and Texas City) discussed in the text.

CH₂O precursors) in the Houston atmosphere? (2) Can the fast CH₂O measurements reported here be used to accurately extrapolate intermittent, lower resolution CH₂O precursor measurements made on the same aircraft? (3) To what extent did the important CH₂O precursors contribute to rapid ozone production in the Houston area? and (4) Were emission inventories accurate with respect to the important CH₂O precursors?

2. Experimental

2.1. Airborne Sampling

[8] The measurements considered here were made from the NCAR L-188C Electra aircraft. CH₂O measurements were made by the NCAR tunable diode laser absorption spectrometer (TDLAS). Supporting measurements included 1 Hz ozone, CO [Holloway *et al.*, 2000], CO₂, NO, NO₂, NO_y [Ryerson *et al.*, 2000], HNO₃ [Huey *et al.*, 1998; Neuman *et al.*, 2002], SO₂, size-resolved aerosol number [Brock *et al.*, 2000], and spectrally resolved actinic flux. Peroxyacyl nitrate compounds were measured once every 3.5 min by gas chromatography (GC) (F. Flocke *et al.*, The behavior of PANs and the budget of total reactive nitrogen (NO_y) during the TOPSE campaign, manuscript in preparation, 2002). Acetaldehyde was measured by proton transfer mass spectrometry (PTRMS) [Lindinger *et al.*, 1998]. Thirty-nine whole air canister samples, typically integrated over 8 s, were acquired on each flight at the discretion of the flight scientist; subsequent GC analysis provided an extensive suite of speciated VOCs (see Table 1) [Schauffler *et al.*,

Table 1. VOCs Measured by the NCAR Whole Air Sampler (WAS)

Alkanes	Alkenes	Aromatics	Alkynes
methane	ethene	benzene	ethyne
ethane	propene	toluene	
propane	1-butene	ethyl benzene	
isobutene	isobutene	<i>m</i> -xylene, <i>p</i> -xylene	
<i>n</i> -butane	<i>t</i> -2-butene	<i>o</i> -xylene	
isopentane	<i>c</i> -2-butene	styrene	
<i>n</i> -pentane	1, 3-butadiene	isopropyl benzene	
<i>n</i> -hexane	3-methyl-1-butene	<i>n</i> -propyl benzene	
2, 2-dimethyl butane	1-pentene	1, 2, 3-trimethyl benzene	
2, 3-dimethyl butane	isoprene	1, 2, 4-trimethyl benzene	
3-methyl pentane	<i>t</i> -2-pentene	1, 3, 5-trimethyl benzene	
2-methyl hexane	<i>c</i> -2-pentene		
3-methyl hexane	2-methyl-2-butene		
2, 2, 4-trimethyl pentane	cyclopentene		
<i>n</i> -heptane	α -pinene		
<i>n</i> -octane			
<i>n</i> -nonane			
<i>n</i> -decane			
cyclopentane			
cyclohexane			
dimethyl cyclopentane			
dimethyl cyclohexane			

1999]. In situ GC samples were acquired once every 15 min, measuring a different subset of VOCs, including a limited suite of oxygenates [Goldan *et al.*, 2000].

[9] Flights surveyed the greater Houston metropolitan area on 8 different days. Two days each were dedicated to exploring the Dallas–Fort Worth area, and power plant plumes throughout eastern Texas. Flight tracks usually involved multiple crosswind intercepts of one or several plumes at a number of distances downwind, in order to characterize the temporal evolution of the plume chemistry. The bulk of sampled plumes were either of mobile source (urban), petrochemical facility, or power plant origin, with many plumes representing some combination of these emission types. In order to capture midday photochemistry and the daily ozone maximum, flights typically took off between 1000 and 1100 CST, and lasted for about 6 hours.

2.2. TDLAS System

[10] The TDLAS and its forerunner systems have been described [Fried *et al.*, 1998, 2002] (B. P. Wert *et al.*, Design and performance of a tunable diode laser absorption spectrometer for airborne formaldehyde measurements, submitted to *Journal of Geophysical Research*, 2002, hereinafter referred to as Wert *et al.*, submitted manuscript, 2002), thus only a short summary is included here. Gas-phase CH₂O was measured by probing the 2831.6417 cm⁻¹ rovibrational line with a tunable diode laser coupled to a 100 m path length astigmatic Herriott absorption cell. Extensive field and laboratory analysis has shown that the TDLAS inlet and Herriott cell produce only minimal sampling artifacts under the conditions encountered during TexAQS [Wert *et al.*, 2002]. Frequent instrument zeros were acquired, and flowing gas-phase standards were generated via a permeation oven, the accuracy of which has been

determined to be $\pm 6\%$ [Gilpin *et al.*, 1997; Fried *et al.*, 1998, 2002]. During the first portion of the TexAQS intensive, ambient CH₂O measurements were made at 10 s resolution. Data integration was reduced to 1 s beginning on the 1 September flight, and this resolution was maintained for the remainder of the mission. Estimated TDLAS precision during the TexAQS mission was 120 parts per trillion by volume (pptv) for 10 s measurements, and 400 pptv for 1 s measurements (reported at the 2 σ level) (Wert *et al.*, submitted manuscript, 2002). Ambient measurements covered roughly 50% of the total flight time, excluding takeoff and landing.

[11] Extensive tests have found that only methanol (CH₃OH) interferes with this technique. Laboratory studies using the same TDLAS arrangement as in TexAQS have shown the methanol interference to be +3.8% at equal methanol and CH₂O concentrations [Fried *et al.*, 2002] (Wert *et al.*, submitted manuscript, 2002), consistent with independent measurements made at Rice University [Richter, 2001]. Thus ambient methanol levels of 10 parts per billion by volume (ppbv) would cause the TDLAS CH₂O measurement to be too high by 380 pptv. TDLAS CH₂O data were corrected using an algorithm derived from 144 coincident measurements of methanol (in situ GC) and CH₂O made on the Electra. Given the 1–3 ppbv CH₂O boundary layer background in the Houston area, and the emphasis of the present study on plume chemistry, the methanol bias and associated correction were largely negligible for the data of interest (<10%). All analyses in this paper utilize methanol corrected data.

[12] Further evidence for the accuracy of the TDLAS CH₂O measurement was provided by comparison with UCLA's ground-based differential optical absorption spectrometer (DOAS) [Stutz and Platt, 1997; Harder *et al.*, 1997] during TexAQS; these two systems were also successfully compared throughout the SOS 99 campaign (Wert *et al.*, submitted manuscript, 2002). During TexAQS the Electra aircraft conducted overflights of the La Porte airport, the site of extensive chemical instrumentation including the UCLA DOAS. TDLAS data was averaged for a 5–10 s period of closest approach in all but one case, corresponding to a path comparable in length (0.5–1 km) and height (40–60 m flight altitude) to that of the DOAS (1.9 km path length, 2–50 m height above ground). Ground-based O₃ and CO measurement variability was acceptably small during the DOAS integration periods (20 min to 1 hour). Bivariate regression of the TDLAS and DOAS CH₂O data produced the relationship: (TDLAS CH₂O) = 0.95 * (DOAS CH₂O) – 0.70 ppbv ($n = 6$, $\chi^2 = 5.5$), with slope and intercept uncertainties of 0.15 and 0.65, respectively (see Figure 2). CH₂O measurements made by three separate DOAS paths at La Porte (0.75, 1.9, and 5.5 km path lengths) agreed well during all overflight periods but one, indicating good mixing. The overflight associated with conditions of poor mixing was that which yielded the high concentration point. Removal of this point produces a slope of 1.23 (± 0.29). These results suggest that systematic measurement errors associated with the two different techniques are small (< $\approx 10\%$).

2.3. Photochemical Model

[13] The photochemical model used in the current study is similar to earlier plume simulations by, for example, Hov

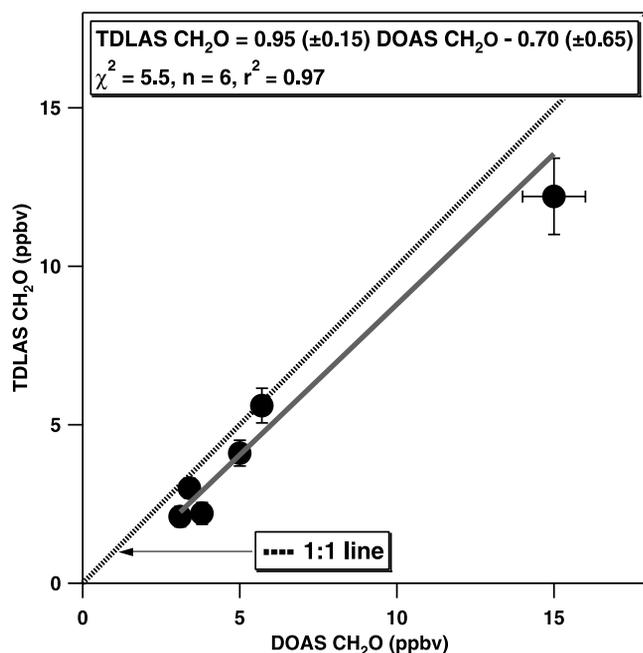


Figure 2. Bivariate regression of CH₂O measurements made by the NCAR TDLAS on the Electra aircraft and the UCLA DOAS at the La Porte site during close overflights.

and Isaksen [1981] and Sillman [2000], where the plume of a power plant was followed in an idealized Lagrangian fashion and the vertical and horizontal spreading were explicitly described. In the present model the vertical dispersion of the plume was controlled by the growth of the planetary boundary layer (PBL), and diffusive vertical transport was calculated as described by Trainer *et al.* [1991]. The maximum height of the PBL was scaled to approximate observed conditions. Horizontal spreading of the plume was described by a diffusion coefficient chosen to approximate the observed growth in plume width as a function of distance. The advective wind was a free parameter that controlled the peak concentration in the plume, and was chosen according to measurements aboard the aircraft, or radar profile measurements made from the ground. Horizontal grid resolution was fixed at 500 m and the plume was emitted into a 100 × 500 × 500 m grid cell. Plume spreading and photochemistry were concurrently calculated within this grid framework. The chemistry was based on the studies of Trainer *et al.* [1987, 1991] where the reaction rates and products were updated according to JPL97 [DeMore *et al.*, 1997] and to the study of Atkinson [1997]. Photolysis rates were calculated by the TUV radiative transfer model of Madronich [1987] and validated by comparison with actinic flux measurements made on the Electra. The TUV model is based upon the discrete ordinates model modified to solve the radiative transfer equation in pseudospherical coordinates [Dahlback and Starnes, 1991].

3. Results

3.1. TexAQS CH₂O Distributions

[14] As depicted in Figure 3, the vertical distribution of CH₂O measured during TexAQS shows a rapid decay above

the boundary layer (0.5–2.5 km) (C. Senff, private communication, 2001), indicative of surface CH₂O precursor sources and the short lifetime of CH₂O under summer conditions. The regional distribution of CH₂O maximums observed below 1000 m flight altitude is shown in Figure 1, at a grid cell resolution of 0.2° latitude by 0.2° longitude (22 × 19 km, respectively). Of primary interest to the present study are the high CH₂O levels found in the greater Houston metropolitan area on many days. These large mixing ratios of between 10 and 32 ppbv were observed in plumes substantially smaller in spatial extent than the Houston metropolitan area. In contrast to Houston, Dallas–Fort Worth had CH₂O levels that were much lower (2–6 ppbv) and more spatially uniform. Measurements of the Dallas–Fort Worth plume are compared with those from Houston, as the populations of the cities are similar (Dallas–Fort Worth: Dallas and Tarrant counties, 3.6 million people; Houston: Harris, Fort Bend, Brazoria, and Galveston counties, 4.2 million people) [U.S. Census Bureau, 2000]. The only power plant plumes with CH₂O levels significantly above background were those from the Welsh and Monticello plants (up to 17 ppbv CH₂O over a very limited area), which encountered strong biogenic isoprene emissions (see Figure 1).

3.2. Houston Area CH₂O Precursor Distributions

[15] To provide an assessment of the relative CH₂O production expected from measured VOCs, the present analysis calculates the “CH₂O production potential” of individual VOC samples [Lee *et al.*, 1998b]. For a given VOC mixture, the CH₂O production potential is here expressed as $\Sigma(k_x \text{OH} * [\text{VOC}_x] * (\gamma_{\text{CH}_2\text{O}}(\text{VOC}_x)))$. Note that in this analysis the CH₂O yield term ($\gamma_{\text{CH}_2\text{O}}$) reflects only the first generation products of OH-initiated VOC oxidation; production from moderately lived intermediates such as acetaldehyde (CH₃CHO) is not considered. Molar CH₂O yields used for the most abundant reactive CH₂O precursors were ethene (1.6) [Niki *et al.*, 1981], propene (1.0) [Niki *et al.*, 1978], and isoprene (0.6) [Carter and Atkinson, 1996]. For the alkanes yields of 0.1 to 0.3 were typically used (exceptions include methane (1) and isobutane (0.8)); the yields of alkynes and most aromatics were set to zero [Lee *et al.*, 1998b]. OH reaction rate constants were given by Atkinson [1990, 1997] and DeMore *et al.* [1997].

[16] CH₂O production potentials were calculated for 390 VOC samples collected in canisters aboard the Electra at flight altitudes less than 1000 m during TexAQS, producing the Houston–Galveston area distribution depicted in Figure 4. A distinct maximum in CH₂O production potential is clearly evident in the vicinity of the Houston Ship Channel. Of the VOC samples taken in this area, six yielded production potentials between 13 and 63 s⁻¹. Only two other samples acquired during TexAQS had CH₂O production potentials greater than 10 s⁻¹. One of these samples was taken in a plume downwind of Texas City, an area of concentrated petrochemical industry south of Houston (see Figure 1). The other was measured in a plume intercepted over eastern Texas, from a petrochemical facility that reported the largest ethene emissions in Texas (Eastman Chemical) [TNRCC, 1999]. On average, the terminal alkenes, largely ethene and propene, composed 95% of total

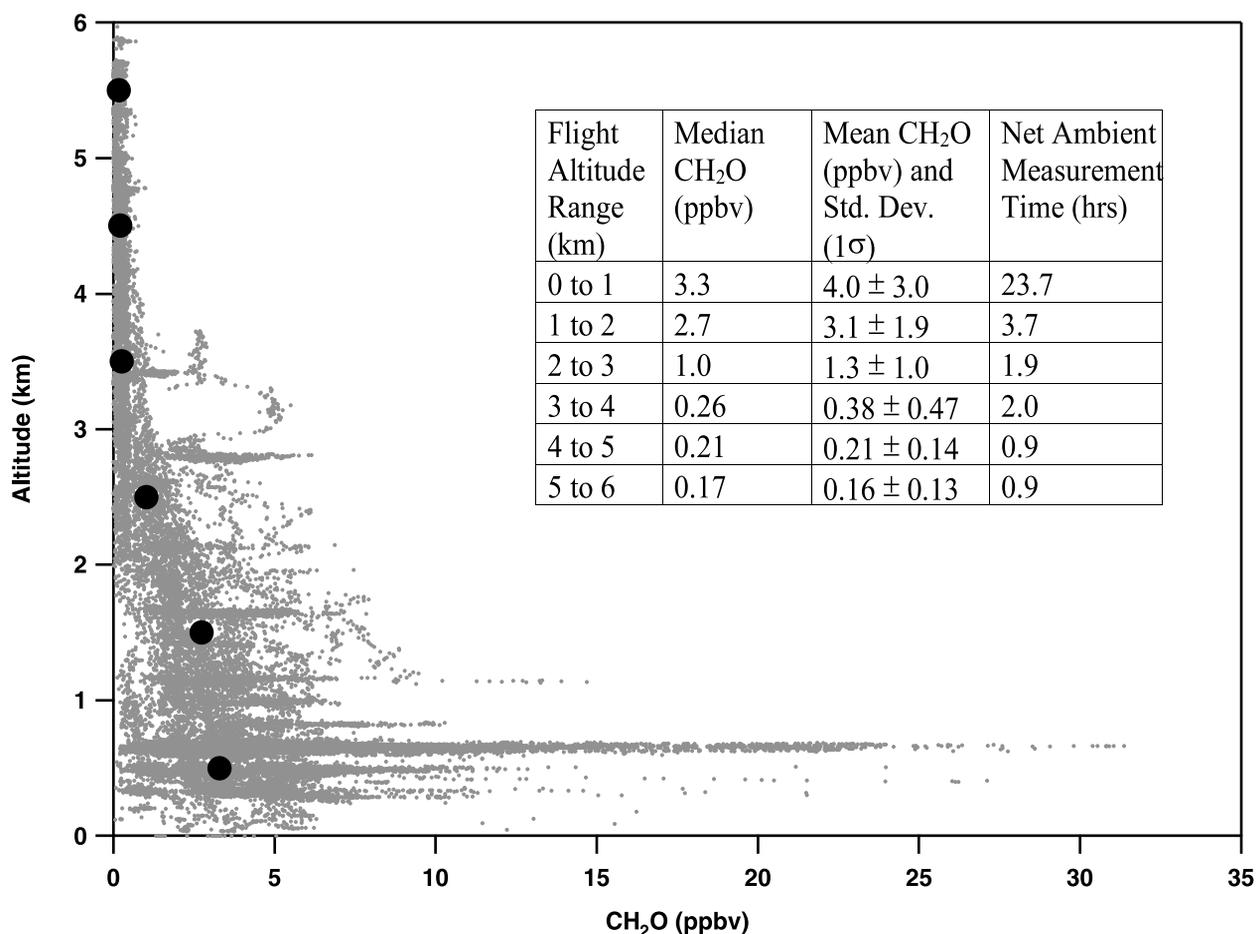


Figure 3. The vertical distribution of CH₂O measurements made during TexAQS 2000. Median CH₂O values for 1000 m flight altitude bins are denoted by the solid circle symbols.

CH₂O production potential and 91% of total hydrocarbon OH reactivity in these eight highly polluted samples. Additional sources of substantial CH₂O production potential were the Sweeny, Freeport, and Chocolate Bayou petrochemical facilities south of Houston (2–9 s⁻¹); these three facilities will be collectively referred to as the “isolated petrochemical facilities.” The forested region northeast of Houston also featured somewhat elevated CH₂O production potentials (1–3 s⁻¹) due largely to isoprene. In all such cases, the bulk of the CH₂O precursors must have been emitted within about an hour prior to sampling, given the rapid midday summertime photochemistry.

[17] The industrial origin of the eight highest CH₂O production potential samples (>10 s⁻¹) was confirmed by ethene/ethyne and propene/ethyne ratios (T. B. Ryerson et al., Effect of co-located petrochemical industrial emissions of reactive alkenes and NO_x on tropospheric ozone formation, in preparation, 2002, hereinafter referred to as Ryerson et al., in preparation, 2002). Strong industrial signatures (ethene/ethyne ≥ 3 and/or propene/ethyne ≥ 1) were identified in 98 canister samples acquired over the greater Houston metropolitan area. These samples had a median ethene/ethyne ratio of 4.9, propene/ethyne ratio of 1.6, CH₂O production potential of 2.9 s⁻¹, and OH hydrocarbon reactivity of 3.1 s⁻¹. The CH₂O production potential of these 98 samples was due largely to the terminal alkenes (78%),

with ethene contributing 30%, propene 22%, and isoprene 14% (1-butene, isobutene, 1, 3 butadiene, 3-methyl-1-butene, and 1-pentene also contributed to a lesser degree). Terminal alkenes were also responsible for a majority (71%) of the total OH hydrocarbon reactivity of these samples.

[18] Seven VOC samples, acquired on 5 different days between 1130 and 1600 CST were selected as representative of Houston mobile source emissions based upon chemical ratios (ethene/ethyne < 1.7 and propene/ethyne < 0.5), CO mixing ratios (>150 ppbv), CO₂ mixing ratios (<375 ppmv), location over the Houston urban core, and wind direction. Median CH₂O production potential (0.6 s⁻¹) and OH hydrocarbon reactivity (0.7 s⁻¹) for these samples were substantially lower than seen in the industrially influenced samples. All nine canister samples acquired over the Dallas–Fort Worth metropolitan area yielded similar signatures, consistent with the lower, more spatially uniform CH₂O levels measured throughout the Dallas–Fort Worth plume (see Figure 1). These observations suggest that the CH₂O species should represent an excellent proxy for the important reactive VOCs in the Houston atmosphere, the utility of which is explored in the remainder of this paper.

3.3. Houston Area Plumes

[19] Measurements of CH₂O in the Houston area are now examined for consistency with the precursor distributions

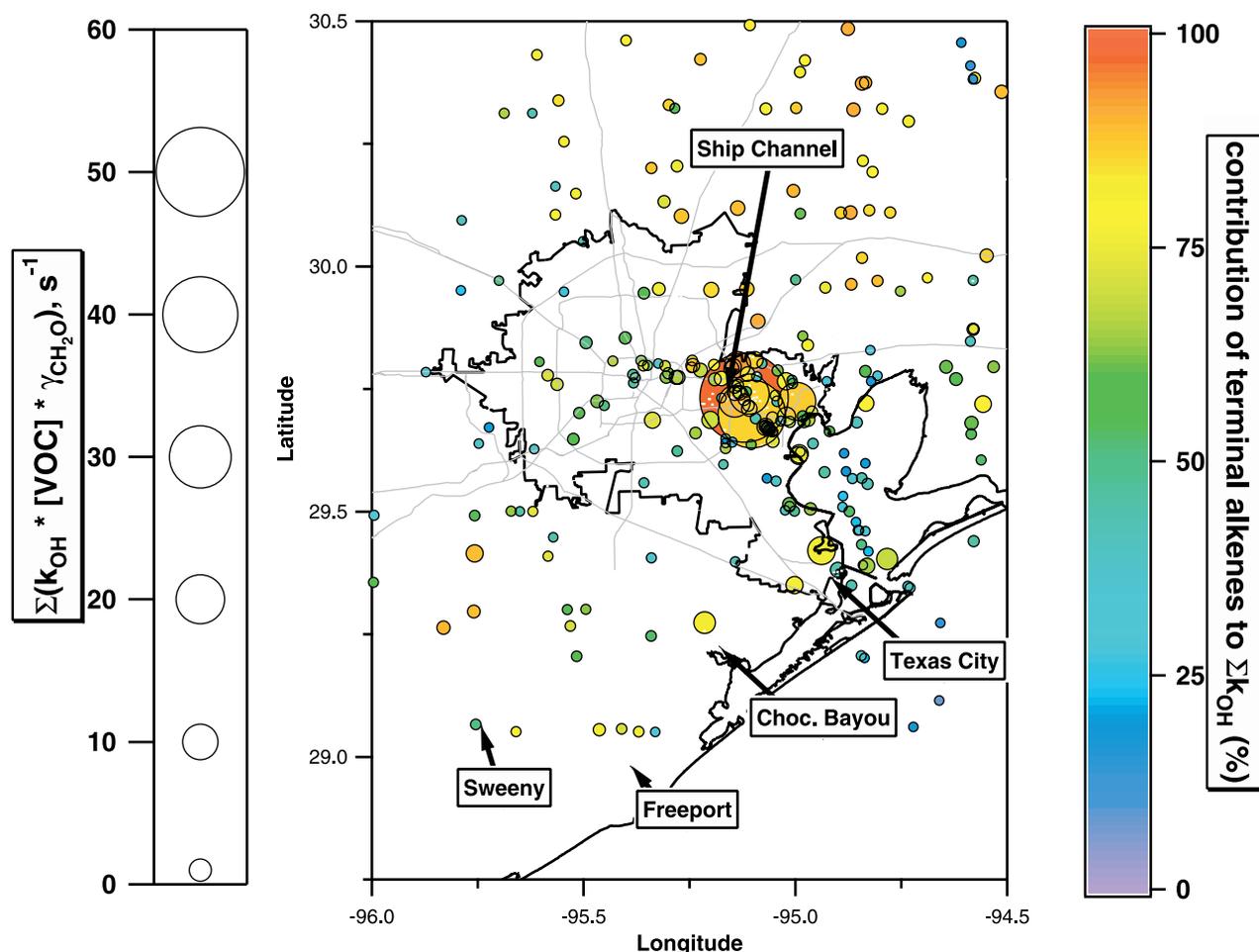


Figure 4. Measured Houston–Galveston VOC distribution coded by size for CH₂O production potential and by color for the percentage of terminal alkene contribution to total OH hydrocarbon reactivity.

discussed above. Complicating this however, lifetimes of ethene (3–8 hours) and propene (1–2.5 hours) are short (lifetime ranges calculated for OH = 1×10^7 mol cm⁻³, O₃ = 120 ppbv, and OH = 4.5×10^6 mol cm⁻³, O₃ = 60 ppbv, respectively) [Atkinson, 1997]. Therefore, regressions of CH₂O versus CH₂O production potentials did not yield strong correlations; this lack of correlation was accentuated by variable meteorology and limited VOC measurements.

[20] Instead, levels of CH₂O in Houston area plumes are compared for transects of roughly equivalent photochemical age. Plume sources were determined using tracers including SO₂, CO, and CO₂ [Brock *et al.*, 2003; Ryerson *et al.*, manuscript in preparation, 2002]. Based on these tracers, distinct power plant (W.A. Parish), mobile source (Houston urban), and petrochemical (Ship Channel) dominated plumes were identified in the 28 August, 30.1° latitude E–W flight transect appearing in Figures 5 and 6. These plumes had passed over the core of the Houston metropolitan area 2.5–3 hours (40 km) previously. Peak NO_y levels in each of the three plumes were roughly comparable (9–14 ppbv), and consistent NO_x/NO_y ratios (0.21–0.27) indicated that the three plumes were well processed and of similar photochemical age. As shown in Figure 6, levels of CH₂O in the petrochemically dominated plume were much higher (up to

13 ppbv) than those of the mobile source (up to 4.5 ppbv) and power plant plumes (up to 3 ppbv).

[21] Within individual plumes transects, tight positive correlations were found between CH₂O and directly emitted species (CO, CO₂, and NO_y), as well as secondary photoproducts (ozone, HNO₃, and acetaldehyde), as listed in Table 2 for two petrochemically dominated plumes. Figure 7 depicts such relationships, showing CH₂O and coincidentally measured NO_y and ozone from a Ship Channel dominated plume transect. Associated correlation coefficient (*r*²) values were greater than 0.95 for the 20-km-wide plume cross section, indicating that the relationship of CH₂O, NO_y, and ozone was very uniform throughout the plume (see Table 2). Note that these good correlations are largely the result of mixing/dilution, as will be discussed below.

[22] When comparing plumes of different sources, large differences were observed in such CH₂O/“tracer” relationships. This is clearly evident in the survey of petrochemical, mobile source (urban), and power plant plumes presented in Table 3. CH₂O/NO_y slopes for petrochemical plumes, including the Ship Channel plume (see Figure 7b), were between 0.8 and 3.1. In contrast, CH₂O/NO_y slopes ranged from 0.2 to 0.4 for the Houston and Dallas–Fort Worth

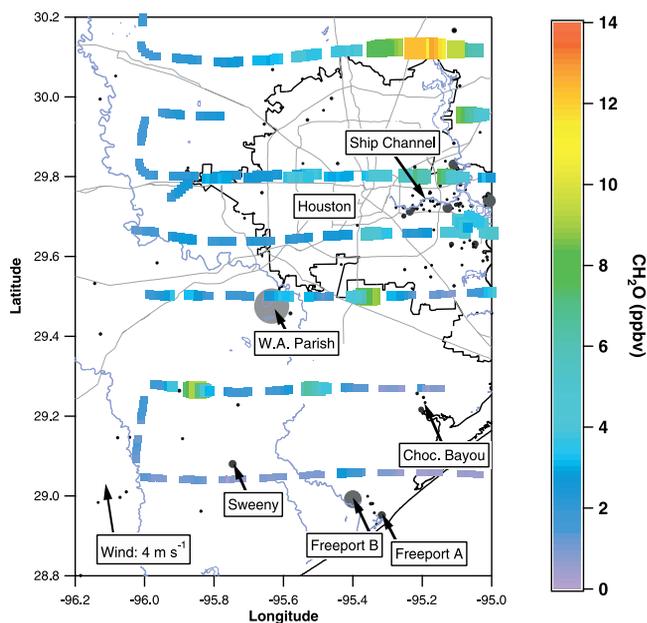


Figure 5. Measurements of CH₂O in the Houston area on 28 August below 1000 m flight altitude. The solid circles show the location of NO_x emitters and are sized according to reported NO_x emissions. General wind direction is shown by the arrow in the bottom left corner.

plumes (mobile source dominated), and the W. A. Parish plume (power plant). Reported CH₂O/tracer slopes were determined from individual plume transects where NO_x/NO_y ratios were generally between 0.2 and 0.3. Some uncertainty in this comparison arises due to HNO₃ loss [Trainer *et al.*, 2000], and the relatively short lifetime of CH₂O that will cause such ratios to change significantly with the photochemical age of the air mass (see Table 2). Nonetheless, the striking and consistent CH₂O/NO_y slope differences found between plumes of different source types on multiple days provides further evidence that petrochemical emissions are indeed the dominant source of extreme CH₂O concentrations in Houston.

[23] The tightly positively correlated CH₂O/O₃ slopes repeatedly measured in mobile source, power plant, and petrochemical plumes (see Tables 2 and 3 and Figure 7c), suggest that oxidation of CH₂O precursors (including photolysis of CH₂O itself) played a major role in ozone production. This is consistent with the observation that CH₂O precursors were responsible for a substantial fraction of total OH hydrocarbon reactivity, especially in the most polluted VOC samples. As depicted in Figure 6, CH₂O and ozone levels observed the same basic trends in the 28 August Houston plumes. Measurements from 1 September, a highly polluted day (see Figures 7a and 7c and 8a and 8b), also show CH₂O and ozone tracking one another. It should be noted that a spatial (temporal) offset was observed between the CH₂O and ozone gradients, as expected, due to the longer lifetime and slower relative net production rate of ozone. Throughout the TexAQS campaign, high CH₂O levels, often 10–20+ ppbv above background, were always found coincident with exceedance levels of ozone (>120 ppbv).

3.4. Isolated Petrochemical Facility Plumes

[24] Using the model discussed above, we now analyze several petrochemical plumes to determine if measured ethene and propene emissions can account for the majority of observed CH₂O and ozone, and if measurements are consistent with NO_x and VOC inventories. For this exercise, plume measurements from the isolated petrochemical facilities, Freeport, Sweeny, and Chocolate Bayou are used [Brock *et al.*, 2003; Ryerson *et al.*, manuscript in preparation, 2002]. These facilities primarily conducted olefin-producing processes, and are analogous to many of the larger petrochemical facilities located along the Houston Ship Channel (C. Wiedinmyer and G. Frost, personal communication, 2001). On 27 and 28 August, plumes from these facilities were superimposed on clean background air, the result of prevailing winds from the Gulf of Mexico (see Figure 5). This kept the plumes largely unperturbed by other emission sources for several hours. Plume model results presented here are for 28 August, as intermittent CH₂O instrument problems were encountered on 27 August. Sufficient CH₂O data were collected on 27 August to demonstrate that the Freeport and Sweeny plumes exhibited consistent behavior on both days. Emission estimates inferred from aircraft observations are reported for Freeport and Sweeny, but not Chocolate Bayou, as VOC data were collected only on the very edge of the Chocolate Bayou plume. Measurements of acetaldehyde and CH₂O were used to apportion emissions of ethene and propene, since propene oxidation produces acetaldehyde while ethene oxidation does not [Niki *et al.*, 1978, 1981; Grosjean and Grosjean, 1996; Grosjean *et al.*, 1996; Calvert *et al.*, 2000; Bacher *et al.*, 2001].

[25] Model NO_x and VOC inputs were derived from observations made 1–2 hours downwind of the petrochemical complexes on both 27 and 28 August, between about 1200 and 1300 CST. It was assumed that by this time the boundary layer was well mixed vertically and the measurements at the flight altitude of 500–600 m were representative of the whole boundary layer (species ratios were constant). Uncertainty in estimated NO_x emission rates for the isolated petrochemical facilities was largely due to model parameters such as wind speed (± 1 m s⁻¹), and boundary layer heights. Considering these uncertainties, the adopted and reported NO_x emission rates for the isolated petrochemical complexes compared favorably (see Table 4). Previous plume studies provide precedent for this good agreement [Ryerson *et al.*, 1998].

[26] Estimates of VOC emissions were derived using measured plume VOC/NO_x ratios, with corrections applied for VOC and NO_x losses occurring between the point of emission and measurement. For the Freeport plume, six separate VOC measurements were made at a variety of downwind distances. One VOC sample piece was collected from the Sweeny and Chocolate Bayou plumes. Due to the dominance of ethene and propene in these samples, the model assumed that VOC emissions for each facility were composed exclusively of these two terminal alkenes. Isoprene was also measured in all plumes, the majority of which was attributed to biogenic emissions from the discontinuous forests of the region, based upon VOC samples taken outside of the plumes and biogenic inventories [Wiedinmyer *et al.*, 2001]. A background isoprene flux

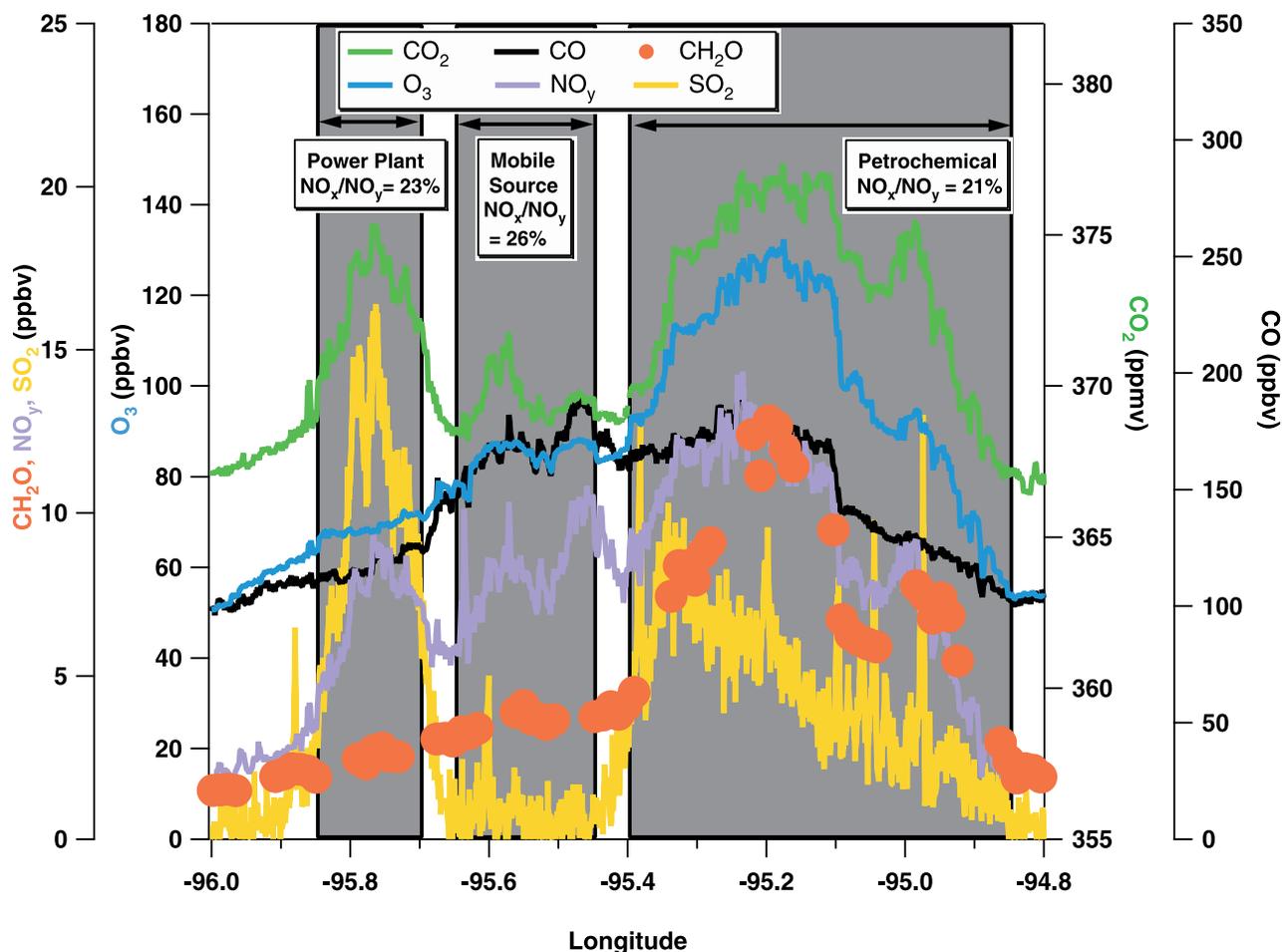


Figure 6. Mixing ratios of CH₂O and other species in a cross section of the Houston plume acquired at 30.1° latitude on 28 August, showing distinct power plant (W. A. Parish), mobile source (Houston urban), and petrochemical (Ship Channel) dominated plumes.

common to all of the plumes was therefore included in the model ($5.2 \times 10^{15} \text{ mol m}^{-2} \text{ s}^{-1}$).

[27] Using these observationally derived NO_x and VOC emission values, the model accurately replicated measured ethene, propene, CH₂O, acetaldehyde, (NO_y – NO_x), (NO_y – (NO_x + HNO₃)), and ozone concentrations. As noted above, observationally derived NO_x values were effectively equivalent to inventory values. Figure 9 provides an overview of results for the 29.3° latitude intercept of the Sweeny and Freeport plumes. Not only did the model successfully simulate the rapid CH₂O enhancement in the plume, it also accurately captured the CH₂O background due to the area-wide isoprene emissions (see Figure 9b). Ethene and propene emission apportionment was corroborated by agreement with both CH₂O and acetaldehyde (Figures 9b and 9c). It is important to note that in all cases the model needed only standard OH-initiated VOC oxidation to achieve good agreement with observations. The observed “shoulder” feature seen on the east side of the Freeport plume, particularly apparent in the ozone measurement (Figure 9d), was due to incursion of the Freeport A complex plume into the Freeport B complex plume. The model only attempted to replicate the larger Freeport B plume (see Figure 5).

[28] The Freeport and Sweeny plumes were also modeled using NO_x and VOC emissions taken directly from TNRCC inventories, along with the measurement-inferred biogenic isoprene background. Note that the only real difference between this and the previously discussed model case was the magnitude of the industrial VOC inputs (see Table 4). When using inventory inputs, less than 10% of the measured CH₂O, and 40% of the measured ozone enhancement over background levels were predicted (see Figures 9b and 9d, dotted gray lines). In this case, modeled CH₂O and ozone production was largely due to the presence of the isoprene background, with TNRCC inventory VOCs contributing minimally. On a related note, inclusion of a large measured isobutane emission to the ethene, propene, and isoprene mix used to model the Chocolate Bayou plume (not depicted) produced very little additional CH₂O and ozone, due to the much lower reactivity of alkanes.

[29] One potential weakness of the above analysis was its reliance on just a handful of canister measurements to characterize the VOC emissions of the isolated petrochemical facilities. However, since CH₂O is efficiently produced by the dominant reactive VOCs in these plumes, it represents a good surrogate for plume VOCs. The very

Table 2. Slopes of CH₂O Versus Several Other Species Measured in the 1 September Ship Channel Dominated Plume and the 28 August Freeport Plume

Distance (km)	Age (hours)	NO _x /NO _y	CH ₂ O/NO _y	r ²	CH ₂ O/CO	r ²	CH ₂ O/O ₃	r ²	CH ₂ O/HNO ₃	r ²
<i>1 September Ship Channel Dominated Plume</i>										
10	0.8	0.69	0.47	0.73	0.21	0.89	0.68	0.87	7.7	0.87
25	2	0.49	0.79	0.99	0.18	0.95	0.41	0.96	5.0	0.94
50	4	0.29	0.73	0.84	0.13	0.92	0.16	0.86	2.5	0.80
75	6	0.24	0.87	0.98	0.09	0.97	0.15	0.98	2.8	0.96
100	8	0.22	0.74	0.89	0.05	0.91	0.12	0.86	1.9	0.77
<i>28 August Freeport Plume</i>										
8	0.6	0.77	0.14	0.90	0.07	0.50	na	na	na	na
33	2.3	0.36	1.35	0.95	0.31	0.94	0.25	0.99	3.7	0.92
60	4.2	0.29	0.87	0.93	0.10	0.84	0.15	0.95	2.2	0.85

robust positive CH₂O/tracer correlations measured in the plumes (see Tables 2 and 3) show that facility emissions were sufficiently collocated, and atmospheric mixing adequately strong, as to rapidly (0.5–1 hour) produce plumes of constant species ratios, despite the fact that the facilities were not perfect point sources. The strong boundary layer mixing necessary to create this measured horizontal uniformity would also be expected to produce

vertically homogeneous species ratios as well. Therefore, plume VOC/NO_x ratios measured at flight altitudes of 500–600 m by the Electra should be representative of the entire boundary layer vertical profile, except for very fresh plumes (<0.5 hour under moderate winds) and perhaps for plumes affected by special meteorological conditions such as coastal effects. Support for this conclusion is provided by airborne LIDAR ozone measure-

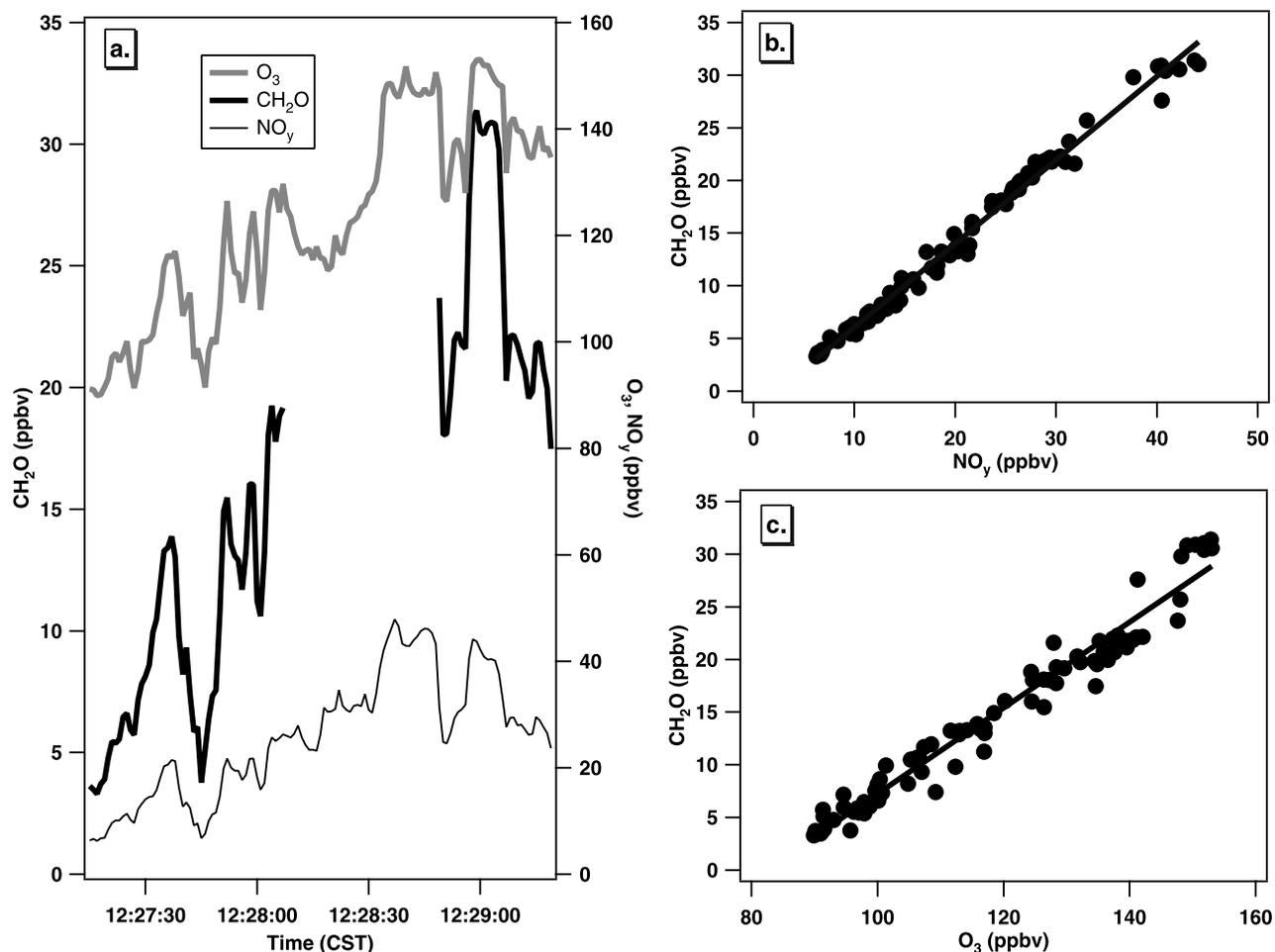

Figure 7. Measurements of CH₂O, NO_y, and ozone in a cross section of the 1 September Ship Channel dominated plume, 2 hours downwind of the Ship Channel (a). Regressions of CH₂O versus NO_y (b) and ozone (c) (see Table 2 for slopes and correlation coefficients).

Table 3. Slopes of CH₂O Versus NO_y and Ozone Measured in Plumes of Petrochemical, Mobile Source (Urban), and Power Plant Origin

Plume Source	Date	NO _x /NO _y	CH ₂ O/NO _y	r ²	CH ₂ O/O ₃	r ²
<i>Petrochemical</i>						
Freeport	27 August	0.37	0.85	0.93	0.19	0.97
Freeport	28 August	0.36	1.35	0.95	0.25	0.99
Freeport	6 September	0.31	1.30	0.78	0.18	0.84
Sweeny	27 August	0.27	1.85	0.92	0.11	0.85
Sweeny	28 August	0.32	1.66	0.99	0.12	0.97
Chocolate Bayou	28 August	0.29	3.13	0.98	0.33	0.99
Ship Channel	27 August	0.24	1.28	0.96	0.15	0.91
Ship Channel	28 August	0.29	0.92	0.82	0.13	0.88
Ship Channel	1 September	0.24	0.87	0.98	0.15	0.98
<i>Mobile Source (Urban)</i>						
Houston	28 August	0.21	0.36	0.81	0.11	0.53
Dallas–Ft. Worth	7 September	0.28	0.25	0.86	0.095	0.80
<i>Power Plant</i>						
W.A. Parish	28 August	0.21	0.24	0.83	0.08	0.94

ments made during TEXAQs that showed strong vertical mixing (C. Senff, private communication, 2001) as well as limited in situ measurement profiles made by the Electra.

[30] Measurement derived VOC emissions for the Freeport and Sweeny plumes used in this model exercise were estimated to be accurate within at least a factor of 2. This was based upon the high degree of agreement achieved between predicted and measured secondary photoproducts, particularly CH₂O and acetaldehyde, and the plume species ratio uniformity demonstrated by CH₂O/tracer relationships. Also accounted for were the uncertainties associated with NO_x emissions, wind speeds (plume ages), and boundary layer heights. Consistent CH₂O/tracer slopes were found on each day the Freeport (3 days) and Sweeny plumes (2 days) were intercepted, as listed in Table 2 for intercepts of comparable photochemical ages (NO_x/NO_y ratios), suggesting that the VOC emissions deduced here were the consequences of standard operations rather than unreported unusual emission events (“upsets”). As seen in Table 4, derived ethene and propene emissions for the Freeport and Sweeny facilities are 50–100 times larger than reported (1999 TNRCC VOC inventory), and in sum about 20 times larger than total reported VOC emissions. Analysis of the Chocolate Bayou plume produced similar disparities.

3.5. Primary CH₂O Emissions

[31] The isolated petrochemical facility plumes were also used to assess the importance of primary petrochemical CH₂O emissions. If primary emissions were large, they would need to be considered in assessments of emissions and photochemistry made using CH₂O data. However, only small enhancements of CH₂O were found in close proximity intercepts of the isolated petrochemical facility plumes, providing an upper limit on direct emissions. In the Freeport plume, CH₂O levels were 1–2 ppbv above background about 20 min downwind (8 km), on both 27 and 28 August (see the 29.1° latitude pass in Figure 5). A similar enhancement was also measured 30 min (7 km) downwind of the Chocolate Bayou facility on 27 August. In all cases, plume CH₂O levels dramatically

increased downwind of these initial intercepts despite dilution, consistent with rapid photochemical production of CH₂O from emitted VOCs, rather than primary emission. In addition, the model achieved good agreement with observations invoking only photochemistry. The dominance of in situ CH₂O production was also supported by extensive, continuous PTRMS measurements of propene, acetaldehyde, and CH₂O made at the La Porte ground site near the Ship Channel (T. Karl et al., The use of proton-transfer mass spectrometry to characterize VOC sources at the La Porte super site during the Texas Air Quality Study 2000, manuscript in preparation, 2002). Direct CH₂O emissions from mobile sources are not expected to have contributed importantly to peak daytime Houston CH₂O abundances either, based upon work by Zweidinger et al. [1988], Altshuler [1993], and others, and the much lower measured CH₂O levels in the Houston and Dallas–Fort Worth mobile source plumes.

3.6. Photochemistry of a Highly Polluted Plume

[32] Ozone exceedance events measured over Houston by the Electra typically involved ozone enhancements of 60–150+ ppbv above background. In comparison, ozone

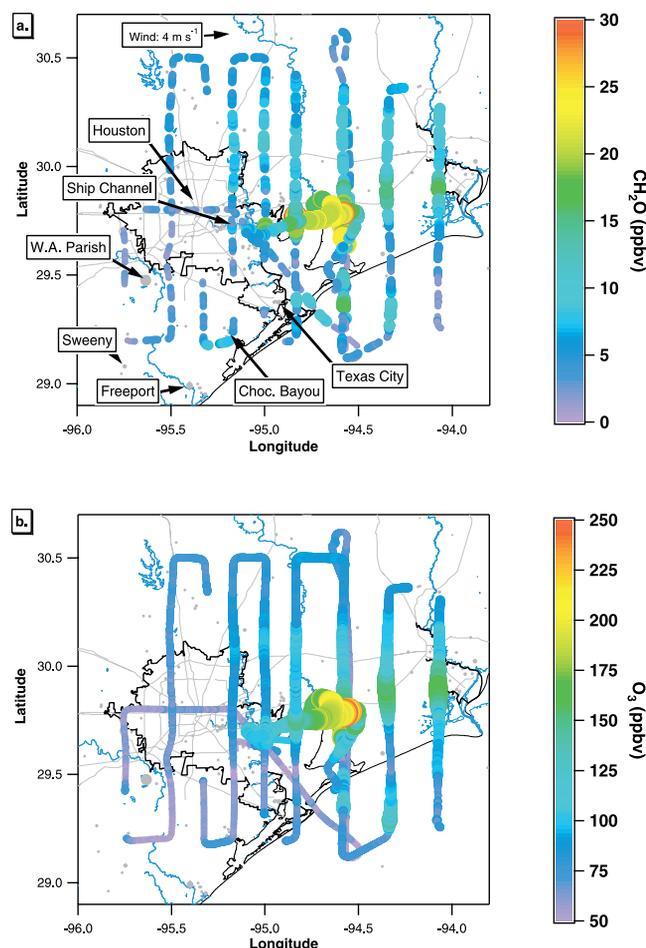


Figure 8. Measured CH₂O (a) and ozone (b) mixing ratios in the Houston–Galveston area on 1 September at flight altitudes less than 1000 m. General wind direction is shown by the arrow in the upper left corner of (a).

Table 4. Measurement-Derived Emissions Used for Modeling the Freeport and Sweeny Plumes (28 August) Along With Corresponding Inventory Values^a

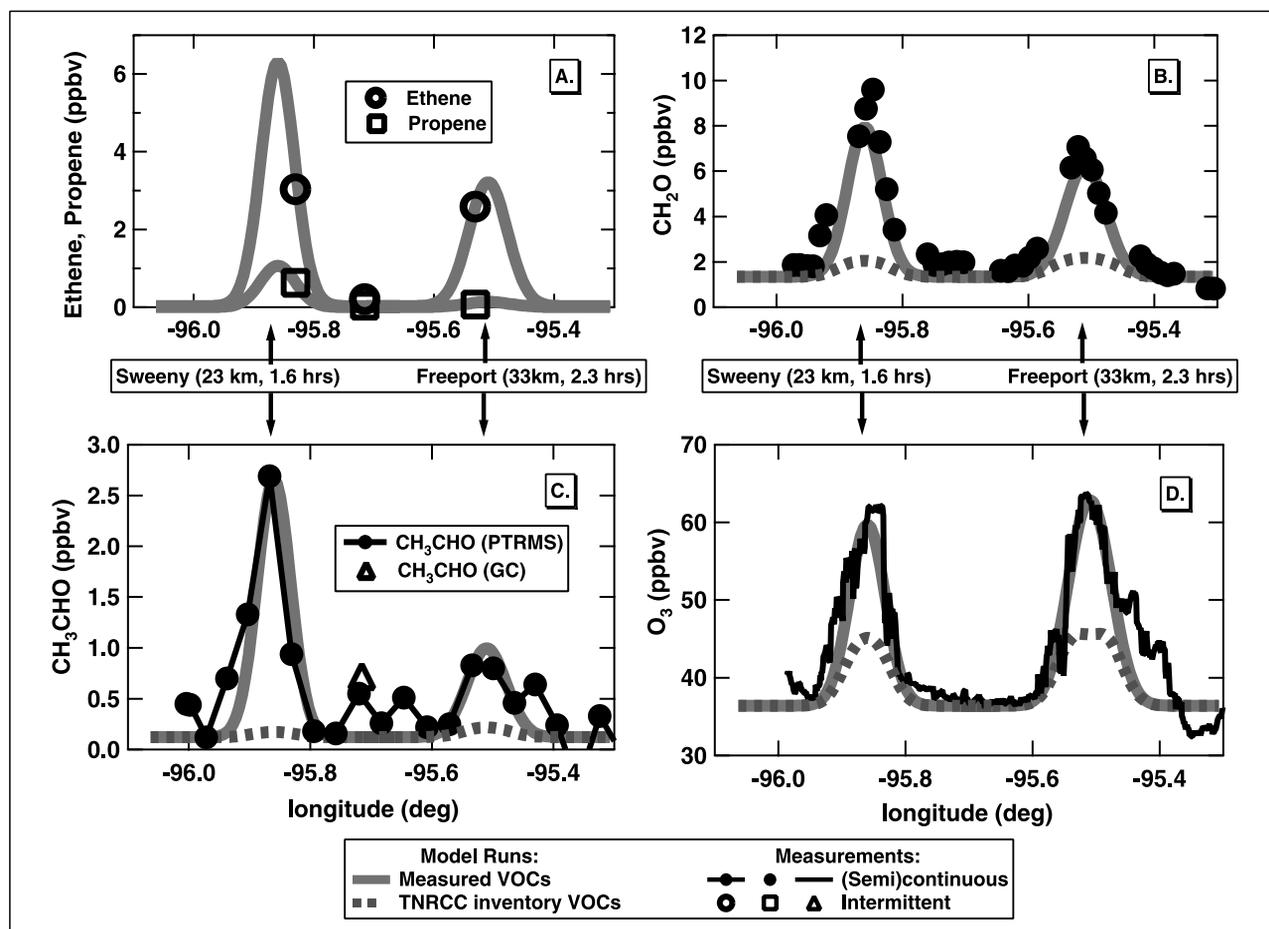
Petrochemical Facility	Measurement-Derived Emissions (kmol hr ⁻¹)			Inventory Emissions (kmol hr ⁻¹)			
	NO _x	Ethene	Propene	NO _x	Ethene	Propene	Total VOCs
Freeport B	30	45	15	31	0.9	0.3	3.1
Sweeny	15	54	30	14	0.5	0.3	3.7

^aInventory numbers from the 1999 and 2000 TNRCC Point Source Database. Freeport B = 6 sources, 28.94° ≤ latitude ≤ 29.02°, -95.44° ≤ longitude ≤ -95.36°; Sweeny = 5 sources, 29.04° ≤ latitude ≤ 29.10°, -95.80° ≤ longitude ≤ -95.7°.

enhancements associated with the isolated petrochemical facility plumes did not exceed 30 ppbv above background. Therefore to test the premise that petrochemical ethene and propene emissions are central to Houston's most extreme ozone pollution, the 1 September Houston plume (see Figures 7 and 8 and Table 2) was modeled using an approach similar to that employed with the isolated petrochemical plumes. This plume was dominated by Ship Channel emissions and had maximum observed ozone levels of 245 ppbv and CH₂O levels of 32 ppbv.

[33] Emission estimates for the 1 September plume were complicated by the diversity and size of the petrochemical source area, and the mobile source emission background,

which although of secondary importance was still substantial. Based on a survey of VOC/NO_x measurements made in the plume, corrected for differential losses, an equal molar emission rate for NO_x, ethene, and propene (252 kmol hr⁻¹ each) was estimated. Nonmethane hydrocarbons (NMHCs) other than ethene and propene contributed only weakly to total hydrocarbon OH reactivity measured in the plume and were left out. The point of this modeling exercise was to assess whether the oxidation of observed ethene and propene could alone account for the general rate and magnitude of extreme CH₂O and ozone production, rather than to provide a detailed check of emission inventories. Use of these simplified inputs yielded good measurement–model


Figure 9. Measurements and corresponding model results for the 29.3° latitude intercept of the Sweeny and Freeport plumes (2.3 hours downwind of Freeport). Shown are ethene and propene (a), CH₂O (b), acetaldehyde (c), and ozone (d) mixing ratios.

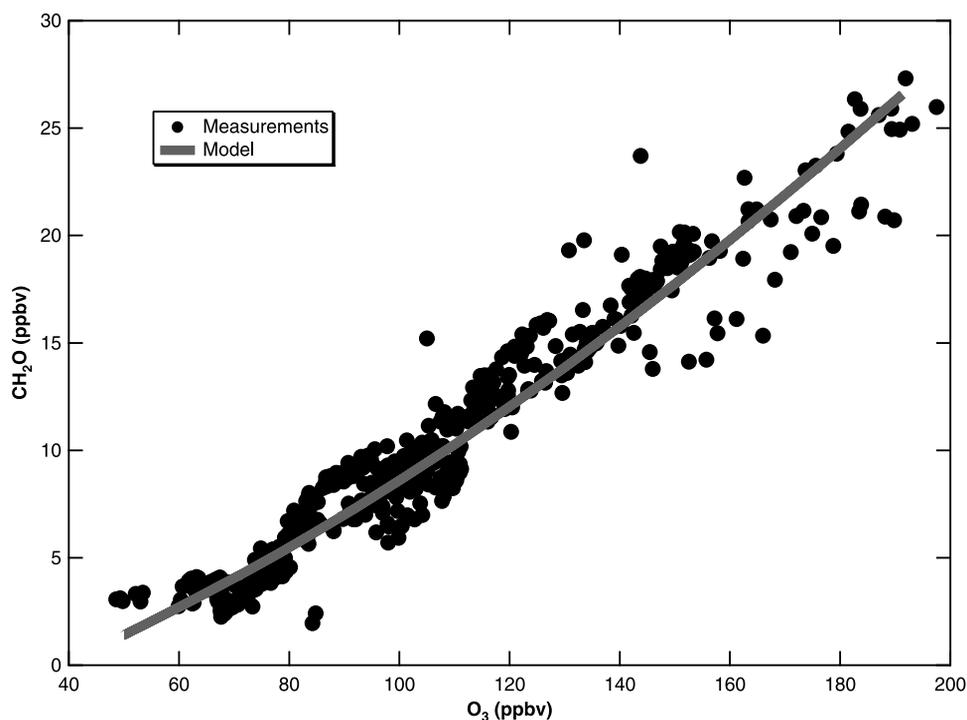


Figure 10. Measured and modeled CH₂O versus ozone for the 1 September intercept of the Ship Channel dominated plume at -94.6° longitude (4 hours downwind of the Ship Channel).

agreement. Figure 10 depicts results for the 94.6° longitude, 4 hours (50 km) downwind intercept (see Table 2), featuring CH₂O concentrations of 3–27 ppbv, and ozone of 50–200 ppbv; NO_y partitioning (not shown) was also accurately replicated.

4. Summary and Conclusions

[34] TDLAS CH₂O measurements made during TexAQS 2000 yielded several important insights concerning terminal alkene emissions and photochemistry in Houston. These insights were amplified by the high resolution (1 and 10 s) of the measurements, the first time such resolution has been achieved in the field for CH₂O.

[35] Results reported here indicate that measured petrochemical ethene and propene levels were alone sufficient to explain the highest CH₂O and ozone levels measured in several Houston area plumes, including levels over 30 and 200 ppbv, respectively. No evidence was found for strong direct emissions of CH₂O relative to secondary production. While enhanced concentrations of VOCs other than ethene and propene were measured in Houston, their lower reactivities rendered them far less important to rapid ozone formation. In addition, the successful modeling indicates that OH-initiated terminal alkene oxidation is the dominant source of ozone in Houston, and that other oxidants such as Cl radicals (and thereby alkanes) play a comparatively modest role, consistent with studies by P. L. Tanaka et al. (Direct evidence for chlorine-enhanced urban ozone formation in Houston, TX, submitted to *Environmental Science and Technology*, 2002).

[36] Modeling of the isolated petrochemical facilities using measurement derived VOC emissions confirms that ethene and propene emissions are far larger than inventories

indicate. Given that these facilities are expected to be representative of the Houston petrochemical industry, and that similar disparities exist between inventory and measured VOC/NO_x ratios for the Ship Channel, it is highly likely that VOC emission inventory deficits are widespread in Houston. Accurate emission inventories will be critical to the success of three-dimensional modeling efforts and the development of efficient ozone control strategies for Houston.

[37] **Acknowledgments.** The TexAQS 2000 mission would not have been possible without the able and diligent efforts of the Southern Oxidants Study (SOS) and Texas National Resource Conservation Commission (TNRCC). We are also very grateful to the National Center for Atmospheric Research (NCAR) Research Aviation Facility (RAF) for providing the Electra aircraft, accomplished on very short notice. NCAR is sponsored by the University Corporation for Atmospheric Research (UCAR). In addition to the National Science Foundation (NSF) Global Troposphere Chemistry Program (GTCP) funds that have supported the long-term development of the NCAR TDLAS system, the work described here was funded in part by grants from the National Oceanic and Atmospheric Administration (NOAA) Climate and Global Change Program and from the SOS program of the U.S. Environmental Protection Agency (EPA). Coauthors Armin Hansel and Armin Wisthaler were supported by NOAA grant NA06GP0483.

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